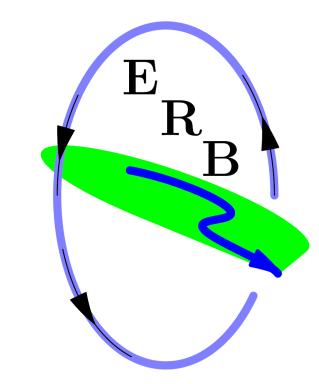


## Lange Bramke



# Bramke basin, Germany

Basin characteristics		ln	Instrumentation and data			
River Basin / River Basin (according EU-WFD) Operation (from to)	Oker river basin / Weser river basin Since 1948, still in operation	Measured hydrological parameters	Measuring period	Temporal resolution	Number of stations	
Gauge coordinates / Gauge datum:	10°26'E; 51°52'N / 537.76 m a.m.s.l.	Stream flow	Nov 1948 – cont.	1h 10 min (since 1992)	1	
Catchment area: Elevation range:	0.76 km² 538 – 700 m a.m.s.l.	Precipitation	1949 – cont. 1980 – cont. 1992 – cont.	Daily Hourly	2	
Basin type: ( alpine, mountainous, lowland) Climatic parameters:	Mountainous 1240 mm (1949-2007), 6.5°C (1962-2007)	Air temp., humidity	1992 – cont. 1987 – cont.	Impuls/ 0.1mm 1h / 10 min.	2	

#### (mean precipitation, temperature and others)

Land use:

Soils: Geology:

Hydrogeology: (Type of aquifers, hydraulic conductivity)

Characteristic water discharges: (Qmin, Qmax, Qmean)

90% Norwegian spruce, 10% grassland

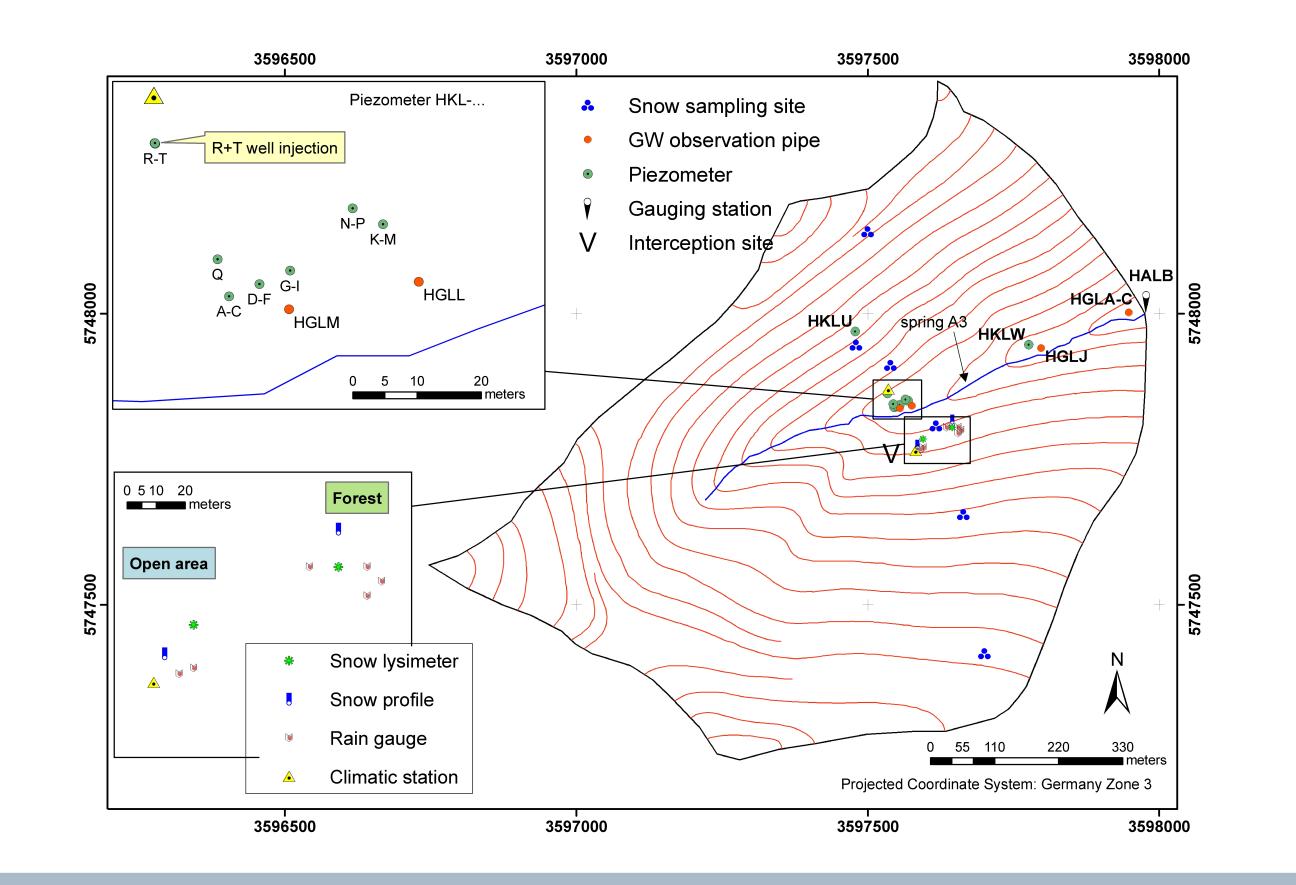
Podsolic brown earth, brown earth Podsol, Pseudogley

Sandstones, shaly quartzite

Fractured rock aquifer with a shallow porous aquifer overlay along the stream channel

0.0 I/s, 15.79 I/s, 634 I/s (1949-2007)

## Map of the research basin

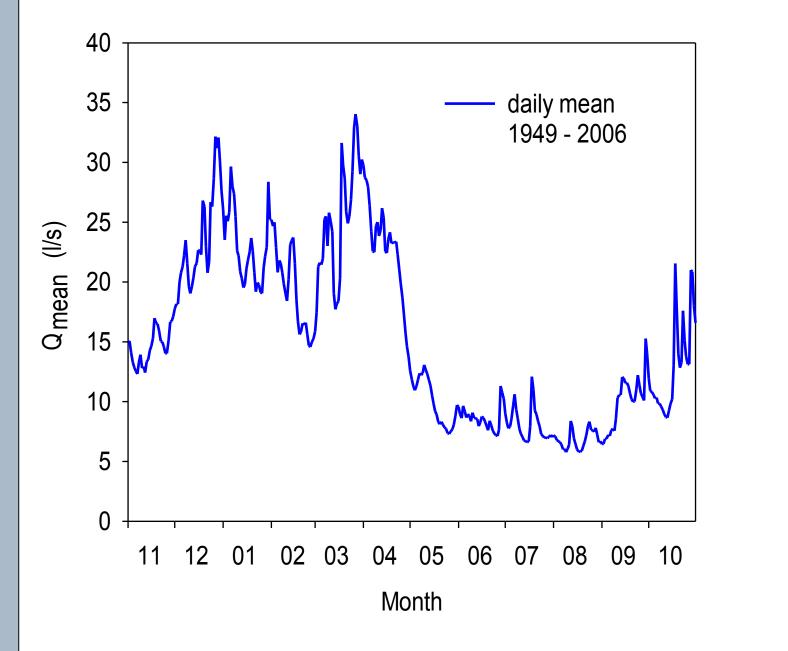


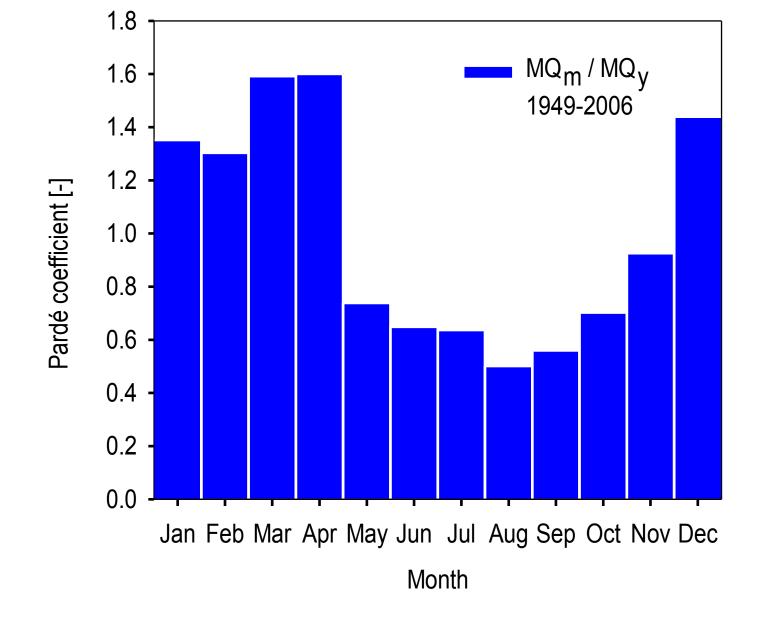
Groundwater level	1988 – cont. 1992 – cont.	Weekly, Hourly	21 5				
Environmental isotopes <sup>3</sup> H, <sup>2</sup> H, <sup>18</sup> O	Event based	Event dependent	Event dependent				
Applied models							
1. Concept model	2. Mike Basin						
3. Feflow	4. IHACRES						

#### Main scientific results

- No overland flow exists. Direct runoff (=event water) equals to 12 %(= 80 mm/a) of total which is less than 5% of the input. Since interflow is negligible, indirect flow is 88 % (= 590 mm/a) and consists of groundwater.
- 2. The unsaturated zone and the fissured rock aquifer are short-cut by preferential flow paths which enable fast percolation. Groundwater recharge is extremely high (620 mm/a). The mean transit time of groundwater is 2.0 years.
- 3. Subsurface pressure heads and groundwater exfiltration react spontaneously on basin input. The latter controls the generation of flood hydrographs which consist dominantly of groundwater. Major cross-faults function as efficient drain channels.

## Mean hydrograph / Pardé flow regime





#### 4. The successive steps of the runoff formation process are probably the following:

(i) Infiltration with saturation of top soils, quick drainage through macropores towards greater depths, and compression of the capillary fringe which may initiate pulse pressure transmission and connected aquifer reactions without mass transfer;

(ii) Rise of piezometric table, i.e. increase of subsurface pressure head and subsequent mass displacement, which can be split into vertical seepage in the unsaturated (cf. (i)) and lateral (groundwater) flow in the saturated zone; and

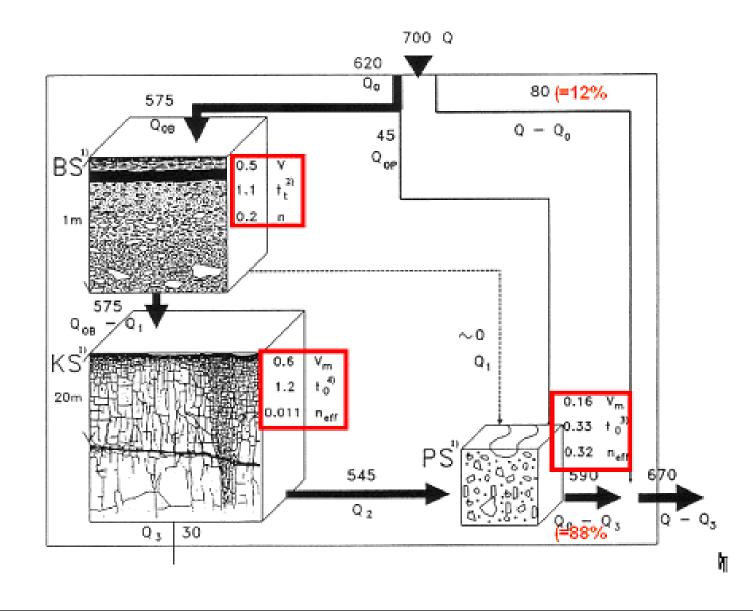
(iii) Groundwater exfiltration to stream channels as a combined effect of pressure transmission and mass transfer, with hydrograph generation as a result. To maintain the quantitative input/output balance, short-term groundwater losses are compensated without much delay, i.e. groundwater recharge is a permanent process throughout the year.

#### Key references for the basin

- Herrmann, A. & Schumann, S. (2009) Untersuchung des Abflussbildungsprozesses als Kontrollmechanismus f
  ür den Gebietswasserumsatz des Oberharzer Einzugsgebiets Lange Bramke (Investigations of the runoff formation process as a mechanism for monitoring the basin turnover in the Lange Bramke catchment, Upper Harz Mountains). Hydrologie und Wasserbewirtschaftung 53(2), 64-79
- 2. Herrmann, A. (2008): 30 Jahre integraler Forschungsansatz zum Abflussbildungsprozess und 60 Jahre

# Special basin characteristics (hydrogeology, lakes, reservoirs etc.)

Mean annual water fluxes [mm WC] and hydraulic reservoir features



 $\begin{array}{l} \mathsf{BS} = \mathsf{Soil} \ \mathsf{water} \ \mathsf{reservoir} \\ \mathsf{KS} = \mathsf{Fractured} \ \mathsf{rock} \ \mathsf{groundwater} \ \mathsf{reservoir} \\ \mathsf{PS} = \mathsf{Porous} \ \mathsf{groundwater} \ \mathsf{reservoir} \\ \mathsf{Q} = \mathsf{Water} \ \mathsf{flux} \ [\mathsf{mm/a}] \\ \mathsf{V}; \ \mathsf{Vm} = \mathsf{Total} \ \mathsf{volume}; \ \mathsf{volume} \ \mathsf{of} \ \mathsf{mobile} \ \mathsf{water} \ [10^6 \ \mathsf{m}^3]; \\ \mathsf{t}_{\mathsf{t}} = \mathsf{Mean} \ \mathsf{transit} \ \mathsf{time} \ \mathsf{of} \ \mathsf{tracer} \ [\mathsf{a}]; \\ \mathsf{t}_{\mathsf{o}} = \mathsf{mean} \ \mathsf{transit} \ \mathsf{time} \ \mathsf{of} \ \mathsf{water} \ [\mathsf{a}] \\ \mathsf{n} = \mathsf{Total} \ \mathsf{porosity} \\ \mathsf{n}_{\mathsf{eff}} = \mathsf{Effective} \ \mathsf{porosity} \end{array}$ 

Abflussbeobachtungen im Oberharz.(30 years of integrated scientific investigation of the runoff formation process and 60 years of runoff observations in the Upper Harz Mountains). Hydrologie und Wasserbewirtschaftung, 52 (3), 132-136

3. Maloszewski, P., Herrmann A., Zuber, A. (1999) Interpretation of tracer tests performed in fractured rock of the Lange Bramke basin, Germany. Hydrogeological Journal, 7: 209-218.

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